

Palaeomagnetic Studies on Rock Formations from Northwest Argentina

K. M. Creer

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III. PALAEOMAGNETIC STUDIES ON ROCK FORMATIONS FROM NORTHWEST ARGENTINA

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Cambrian, Cambro-Ordovician and Ordovician formations of red sandstones and siltstones from the sub-Andean regions of the Provinces of Salta and Jujuy have been studied. The grouping of n.r.m. directions suggests partial remagnetization after folding in the Tertiary or Quaternary geomagnetic field. Thermal cleaning at 300 °C and higher temperatures destroys this secondary magnetization leaving a magnetization which is accepted as representing the lower Palaeozoic geomagnetic field. Palaeomagnetic south pole positions have been computed and lie in the Atlantic Ocean to the NNE of Brazil.

The period between the Carboniferous and Triassic is covered by the Paganzo formation which is exposed in La Rioja Province. The older part (Paganzo II) is reversely magnetized with a south p.m. pole in the South Atlantic, while the upper part (Paganzo III) is normally magnetized with a south palaeomagnetic pole in the vicinity of poles obtained for Triassic rock formations from elsewhere in

Formations of red beds from Salta province regarded as Upper Palaeozoic or Mesozoic yield a south palaeomagnetic pole in the South Atlantic corresponding to a Triassic age.

Ore microscope and thermal decay curves suggest that the remanent magnetism is due to haematite.

1. Introduction

Palaeozoic rocks are found in Argentina principally in the northwestern provinces of Salta and Jujuy and also in the Sierra de la Ventana in the south of the province of Buenos Aires. The latter have been found to be too weakly magnetized to warrant detailed studies. The former consist of red sandstones and siltstones and are highly suitable for palaeomagnetic investigations. In fact it was from these formations that the first reliable estimates of the lower Palaeozoic poles for S. America were obtained and these have been quoted in review papers (Creer 1964, 1965, 1968).

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2. GEOLOGY

The stratigraphy of northwest Argentina is summarized in table 1.

2.1. Pre-Cambrian

In the Pre-Cordillera, the Pre-Cambrian is represented by phyllites, quartzites and crystalline limestones exposed a little to the northwest of Mendoza where they are intruded by gabbros

TABLE 1. STRATIGRAPHIC TABLE, JUJUY

Cenozoic	Quaternary	Recent and Present			sands, clays and pebbles
		Lower	-discordance		conglomerates, sandstones, etc.
			discordance Huachichocana fm		andesite (hornblende)
	Tertiary	Pliocene	-discordance		sandstones and red clays
Mesozoic	Cretaceous		-discordance		limestones, oolitic
Palaeozoic	Ordovician	Arenigian	–discordance–––– Huantes	Cienaguillas fm	lutites and clay- sandstones
		Ur Tremadocian	Group	Coquena fm	sandstones, limestones and lutite
			-discordance		
	Cambrian		Meson	Chalhualmayoc fm	
					ortho quartzites
			Group	Campanario fm Lizoite fm	1
**************************************			-discordance		
Pre-Cambrian			Puncoviscana fm		slates, greywackes and quartzites

and serpentinized periodotites. In the north of Salta Province and in the puna of Jujuy Province green and violet phyllites and greywackes are exposed at the base of the Cambrian over a large area. They are only weakly metamorphosed.

2.2. Cambrian

The geological map of the area makes no distinction between Cambrian and Ordovician formations.

Unfossiliferous Cambrian deposits are known in the eastern Cordilleras of northern Argentina where they rest with strong angular unconformity on highly folded Pre-Cambrian rocks and are covered, with angular unconformity, by the lowermost Ordovician beds. The group attains a maximum thickness of 1500 m and is formed of brightly coloured sandstones and shales. The middle part of the sequence contains frequent intercalations of reddish and green shales and thin quartzitic layers packed with Skolithos tubes.

2.3. Ordovician

The Ordovician has been mapped with the Cambrian in northern Argentina and with the rest of the Palaeozoic in western Argentina. These rocks attain great total thickness and regional extent. They form many subparallel bands of general north-south trend. The Tremadocian is well developed in the eastern Cordilleras of Salta and Jujuy provinces where it attains a maximum thickness of 2400 m. In the western border of the outcrops (Quebrada del Toro, Salta Province) the basal conglomerate grades upward into a green shale and light coloured quartzite with intercalations of glacio-marine conglomerate. These are succeeded by green shales with thin lenticular masses of blue limestone.

In the central outcrops (Quebrada de Humahuaca, Jujuy Province) the lower part of the sequence is formed of black shales, succeeded by sandy limestones and green shales.

In the eastern border (Rio Santa Victoria, Salta Province), the whole sequence is formed of dark grey shale.

Grey-green shales alternating with yellowish sandstones constitute the Arenig in Cerro San Bernado near Salta City.

In the Famatina range of northwest Argentina, the Llanvirnian, consists of 700 m of green and blue shales with dark-grey sandstones capped by 80 to 100 m of reddish shaly sandstone. The beds are unconformably overlain by Mississippian continental strata. The reddish sandstones have yielded fossils. In the western part of the Cordillera, between Calingasta and the Pampa de Yalgueraz, the Caradocian is represented by 800 m of dark greenish greywackes and shales, capped by 200 to 300 m of purple and green phyllitic shales.

2.4. Silurian

In north and west Argentina, Silurian rocks are not differentiated from the rest of the Palaeozoic. In the southern hills of Buenos Aires province they have been grouped with the Devonian.

In the Pre-Cordillera, the Silurian is exposed between Jachal and the San Juan river where it consists of dark greenish grey shales and intercalations of shaly limestone. These beds contain two thick intercalations of highly haematitic layers near the base and rest with slight regional unconformity on the Caradoc.

2.5. Devonian

Devonian rocks are exposed in the sub-Andean ranges of Salta and Jujuy Provinces. The sequence rests with slight unconformity on Silurian strata and attains 800 m. It is formed of dark grey shales with intercalations of grey and pink quartzitic sandstones.

2.6. Carboniferous to Triassic

In the Pampean Ranges at the eastern border of the Precordillera, Mississippian beds rest on peneplained Pre-Cambrian rocks. In the Famatina Range they consist of 1200 m of red, yellow, green, blue and grey sandstones, partly glacial conglomerates, and siltstones and shales (Paganzo I) overlain unconformably by exclusively continental (Paganzo II) beds consisting of yellowish, green, grey and reddish sandstones, shales, and conglomerates which are again in part of glacial origin. The latter are Pennsylvanian and may be up to 700 mm thick.

They are overlain with slight erosional unconformity, in northern La Rioja Province by

the Patquia formation (Paganzo III formation), 1500 m of typical red beds which begin with pink or whitish conglomerates of Permo-Triassic age.

Permian and Triassic beds also occur in the sub-Andean ranges of Salta and Jujuy where they begin with tillites which rest unconformably on the lower Devonian, followed by interbedded green and red shales and 700 m of varicoloured sandstones and shales.

3. PALEOMAGNETIC RESULTS

A sketch map of the area sampled in Salta and Jujuy provinces is shown in figure 1.

3.1. Cambrian and Cambro-Ordovician

Localities and brief descriptions of these samples are listed in tables 2 and 3. Recent geological work allows the outcrops from which samples were collected to be classified as Cambrian or Cambro-Ordovician.

The n.r.m. directions have been plotted on stereographic projections in figure 2 with respect to the present horizontal and in figure 3 with respect to the bedding planes, i.e. the palaeohorizontal. A brief inspection of figure 2 indicates these rocks have been remagnetized in recent geological time.

A representative group (A) of samples whose n.r.m. directions are shown in figure 4 were demagnetized thermally, with results as illustrated in figure 5. These groups of directions both referred to the bedding (i.e. the palaeo-horizontal) show a pronounced improvement after cleaning. The data plotted in figure 5 consist of a population of directions chosen after demagnetization at the optimum temperature for each sample, these temperatures having been chosen by studying the trend in changes of direction and intensity of remanence during cleaning. The sample-mean directions at each demagnetization step are listed in table 4 and the corresponding mean intensities of r.m. are listed in table 5. The optimum demagnetization temperatures are indicated by the entries in italics. Of the 16 samples studied, 478, 550 and 565 were excluded since they appear to be normal while the main group, N=13, are reversed (population A, table 6).

Another set of specimens from these samples was demagnetized in steps. The results are given in table 6 (population B) and the successive mean directions illustrated in figure 7 are seen to be in excellent agreement with the earlier result. The secondary component appears to have blockage temperatures less than 300 °C because the cleaned directions at 300, 500 and 560 °C are closely grouped (figure 7). The palaeomagnetic pole after cleaning at 500 °C has been taken as representative of the primary magnetization of this formation.

The distribution of intensities is log-normal and therefore geometric-mean values have been computed with their appropriate standard deviations and standard errors. The standard errors are illustrated in the thermal decay curve (figure 6), a logarithmic plot having been used in this figure because this makes the error bars symmetrical about the mean values. The blockage temperature of the main component of remanence is distributed over the whole range of temperatures shown: at 500 °C only about 15 % of the natural magnetization remains: between 600 and 660 °C, just over 10 % is left. The remanence measured after cleaning above 570 °C must certainly be due to ferric oxide. That existing below this temperature could be due to either fine grained ferric oxide such that the blockage temperature is less than the Curie temperature or titanomagnetite. The red colour of the beds and ore microscope observations

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described in §4 indicate that ferric oxide is the most likely carrier of remanence in these rocks.

The directions of r.m. of populations A and B after cleaning do not differ significantly at the 95 % probability level, equivalent to twice the standard error (δ_m) circle.

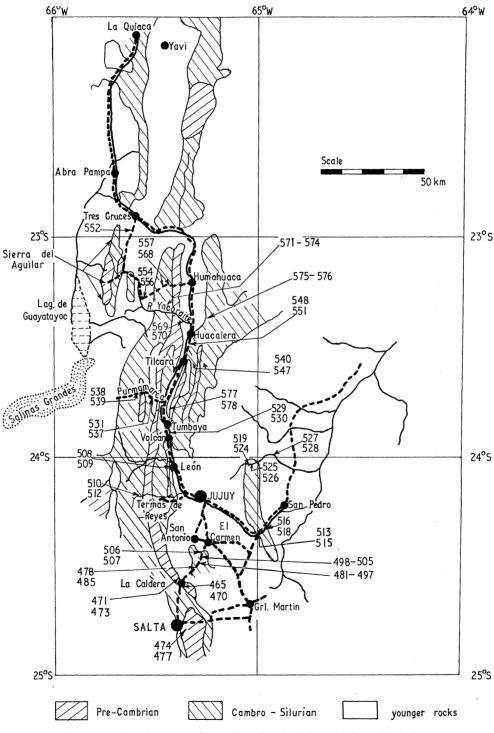


FIGURE 1. Sketch map and sampling sites in Salta and Jujuy Provinces.

Provinces
Jujuy
AND J
SALTA
FROM
CAMBRIAN SAMPLE DESCRIPTIONS FROM SALTA AND JUJUY
SAMPLE
CAMBRIAN
TABLE 2.

	our code no.	pale reddish brown $10R-5/4$ greyish red $10R-4/2$		wn $5 T R - 5/2$ 5 R - 6/2	'dark reddish brown $10R-3/4$ moderate red $5R-5/4$			10R-6/2 $5R-6/2$	greyish red $5R-4/2$ moderate red $5R-5/4$ dark reddish brown $10R-3/4$	$\begin{array}{ccc} {\rm ed} & 5{\rm R}-4/2 \\ {\rm ed} & 10{\rm R}-4/2 \\ {\rm own} & 5{\rm YR}-2/2 \end{array}$	ed $5R-4/2$ ky red $10R-2/2$	s red 5R-5/4 5R-6/2 d 5R-3/4
OVINCES	colour	pale reddisl	greyish red	(pale brown)	(dark reddish)			pale red	greyish red moderate red dark reddish	greyish red greyish red dusky brown	greyish red very dusky red	moderate red pale red dusky red
SALTA AND JUJUY FR	description	v.f.g. sandstone	1	1		1		massive hard m.g. sandstone	massive f.g. sandstone	sheared siltstone	v.f.g. sandstone	f.g. sandstone
KIPTIONS FROM	dip of beds	40° at 260°	18° at 0°	20° at 270°	15° at 240°	25° at 335°		40° at 280°	25° at 285°	95° at 285°	30° at 115°	45° at 275°
TABLE 2. CAMBRIAN SAMPLE DESCRIPTIONS FROM SALTA AND JUJUY FROVINCES	location yuy cities	'Pedrera' road, 18–19 km	'Cornisa' road 28 km, north of Salta	'Cornisa' road, 36.5 km north of Salta	'Cornisa' road, 38 km north of Salta	Yala-Las Lagunas road		Tilcara: 'garganta del diablo' road, 4 km	as site 6 at 4.8 km, older than site 6	as site 6 at 5.9 km, (Lower Cambrian?)	as site 6 at 6 km, across the fault	Perchel
	site no. sample nos. (A) Between Salta and Jujuy cities	478, 479, 480	498–501	502, 503	504, 505	508	(B) North of Jujuy city	540, 541	542–543	544–545	546–547	548–551
	site 1 (A)	-	67	က	4	ĸ	(B)	9	4	∞	6	10

Table 3. Cambro-Ordovician sample descriptions from Jujuy Province

		Tubur of Campio Charles and Campion and Ca				
te no.	te no. sample nos.	location	dip of beds	description	colour	code no.
	506, 507	'Cornisa' road between Salta and Jujuy, 200 m inside Jujuy province	70° at 90°		moderate red	5R-5/4
•	538, 539	Quebrada Purma Marca	95° at 250°	massive siltstones	greyish red very dusky red	5R-4/2 10R-2/2
	557–560	El Aguila-Humahuaca road, 27 km from Humahuaca	30° at 280°	massive f. to m.g. sandstone	(pale red purple pale red light brownish grey very dusky red	5RP-6/2 5R-6/2 5YR-6/1 10R-2/2
	561–562	as site 3, 25.4 km from Humahuaca	30° at 240°	massive f. to m.g. sandstone	greyish red	5R-4/2
	563-564	as site 3, 23.3 km from Humahuaca	40° at 80°	f.g. sandstone	olive grey greyish red	5Y-4/2 $10R-4/2$
	565–566	as site 3, 20.5 km from Humahuaca	30° at 80°	f.g. sandstone	greyish red	5R-4/2
	567-568	as site 3, 17.2 km from Humahuaca	35° at 135°	f.g. sandstone	dark greenish grey	5GY-4/1
	577, 578	Tumbaya	49° at 290°	massive, pisolitic sandstones	-	1

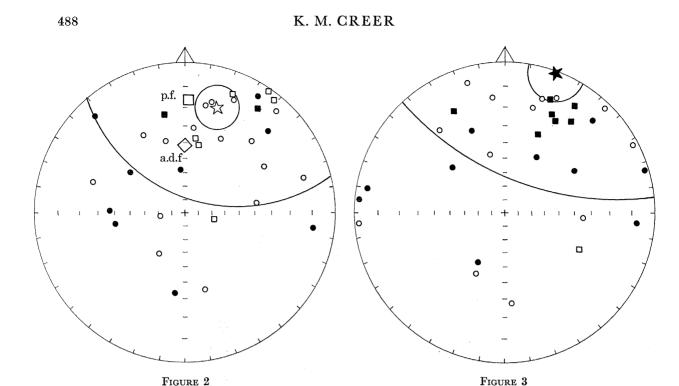


FIGURE 2. N.r.m. directions of whole collection of Cambrian (circles) and Cambro-Ordovician (squares) samples referred to the present horizontal. The mean direction is indicated by the star and the standard error and standard deviation circles are shown. The present field (p.f.) and axial dipole field (a.d.f.) directions are also plotted. Open symbols represent upward inclinations and solid black symbols represent downward inclinations.

FIGURE 3. N.r.m. directions referred to the bedding, i.e. to the palaeohorizontal for the same population of samples as in figure 2.

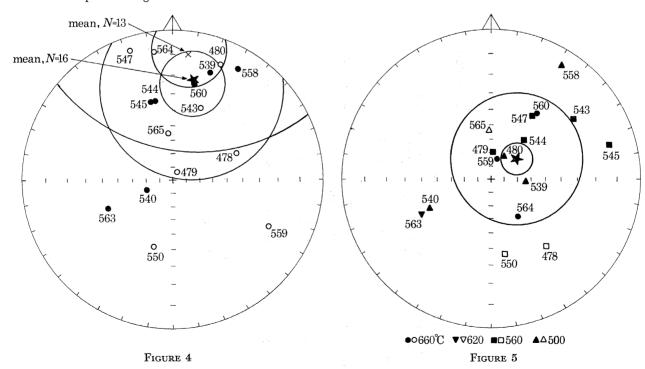


FIGURE 4. N.r.m. directions referred to the bedding of Cambrian samples selected for thermal cleaning. Sample numbers shown.

FIGURE 5. Cleaned directions of r.m. with mean and standard error and standard deviation circles for the population of Cambrian samples of figure 4. Temperature of cleaning indicated by symbol.

Table 4. Thermal demagnetization of Cambrian and Cambro-Ordovician samples

	٥	-27	- 58	+35	8+	ı	+35	+49	+74	+53	- 56	+73	+44	1	+27	+17
	D ₀ 099	331	355	343	162		178	123	309	43	1	20	34	1	142	279
	620 °C	-2 + 25	-25	6-	+53	-1	4	+	9+	99-	+27	+1	*	+35	+ 38	-34
	62(96 272	265	263	217	66	288	108	1117	172	348	320	12	244	342	62
at	260°C	- 29 + 70	1	+64	+43	+55	+53	+12	+37	-39	-37	- 44	+12	1	1	
etization	560	140 3		314	539	23	36	73	30	17.1	58	73	11	1		
demagn	ပ္စ	+ 8 + 81	+71	+65	+40	1	+28	ရ	+37	-32	+	- 19	9+	+43	-21	-53
r thermal	200	123 3	92	94	246	1	13	83	56	185	32	43	12	308	347	357
grees afte	C	$+31 \\ -15$	+43	+41	+30	-64	+47	6+	+26	4	+14	-25	9+	+30	-16	99 –
of r.m./de	400 °C	171	52	34	25I	139	27	79	31	187	53	72	10	248	351	353
directions of r.m./degrees after thermal demagnetization at	ာ ၁	+41 +41	99+	+74	+42	+	+62	+21	+48	- 42	+27	-15	6+		-	-
i ð	D. 008	139 28	52	145	247	6	356	24	61	177	27	86	13	1	İ	I
	ပ္စ္	+ 24 - 80	68 –	+ 60	+40	-52	+52	+28	-1	-45	+15	- 44	+10	+38	-13	- 48
	150	99 + 24 $41 - 80$	38	18	248	15	338	354	342	196	28	97	10	240	351	358
		- 40 - 83	-10	+15	69 +	-35	+33	+33	9-	-41	8+	- 13	+23	+40	81	- 55
	20 °C	68 - 40	23	50	247	22	348	345	342	195	30	114	14	246	352	356
	age	ΦΦ														
	sample no.	478	480	539	540	543	544	545	547	550	558	559	260	563	564	565

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3.2. Ordovician

The sampling sites are illustrated in figure 1 and brief descriptions of the samples are given in table 7. These samples were taken from formations near the top of the Ordovician according to C. M. Turner (1967, personal communication), although in some review and summary articles (Creer 1962, 1964, 1965, 1968) they have been mistakenly quoted as being Devonian, this information having been provided during the field work.

Table 5. Thermal demagnetization of Cambrian and Cambro-Ordovician samples intensity of r.m./uG after thermal demagnetization at

			mu	ensity of r.m	$1./\mu$ G after 1	inermai den	nagnetizatio	n at	
sample									
no.	age	$20~^{\circ}\mathrm{C}$	150 °C	300 °C	400 °C	500 °C	$560~^{\circ}\mathrm{C}$	$620~^{\circ}\mathrm{C}$	$680~^{\circ}\mathrm{C}$
478	\mathbf{e}	2.5	1.1	1.4	1.1	0.8	0.7	10.6	
479	\mathbf{e}	2.7	0.7	1.8	0.2	0.2	0.1	3.4	0.1
480	\mathbf{e}	2.5	50.1	2.1	0.7	1.0		9.9	0.1
539	\mathbf{e}/\mathbf{o}	14.4	5.1	-	2.4	0.5	0.1	49.9	10.8
540	\mathbf{e}	18.3	20.2	13.5	6.6	5.1	11.8	1.9	0.4
543	\mathbf{e}	2.2	1.6	1.9	0.8	-	0.3	5.4	
$\bf 544$	\mathbf{e}	17.1	12.3	4.2	1.9	0.9	0.6	15.8	3.5
545	\mathbf{e}	10.8	4.4	1.6	1.4	1.3	1.0	42.0	3.8
547	\mathbf{e}	19.3	9.7	4.9	1.9	1.8	1.7	30.8	6.4
550	\mathbf{e}	3.4	2.7	1.5	0.9	0.7	0.6	3.8	0.3
558	\mathbf{e}/\mathbf{o}	6.3	5.4	3.9	1.9	1.1	0.4	16.0	0.4
559	\mathbf{e}/\mathbf{o}	2.1	2.4	1.7	1.0	0.6	0.7	0.9	0.3
560	\mathbf{e} /O	95.1	85.1	75.0	44.3	36.9	33.7	17.6	1.1
563	\mathbf{e}/\mathbf{o}	124.6	102.7		16.7	50.1	******	2.3	
$\bf 564$	E/O	18.1	17.1		5.4	7.1	Million Co.	6.0	0.4
565	E/O	8.1	7.5		4.3	3.0		10.9	0.4

TABLE 6. THERMAL DEMAGNETIZATION OF CAMBRIAN AND CAMBRO-ORDOVICIAN

		number	d	irection of	r.m./deg	ree		south pal	aeopole	
population	treatment	of samples	D	I	δ	$\delta_{ m m}$	lat.	long.	δ	$\delta_{ m m}$
A	n.r.m.	13	7	+10	58	16	61° N	57° W	57	16
(see figures 4, 5)	thermally cleaned	13	42	+64	43	12	8° N	40° W	52	14
В	n.r.m.	19	15	+10	68	15	57° N	35° W	65	15
(see figure 7)	150 °C	19	14	+4	73	17	63° N	26° W	71	16
	$300~{\rm ^{\circ}C}$	13	48	+43	66	18	19° N	22° W	68	19
	$500~^{\circ}\mathrm{C}$	18	51	+41	68	16	$15^{\circ}~\mathrm{N}$	22° W	66	16
	$560~^{\circ}\mathrm{C}$	14	59	+43	67	18	8° N	20° W	67	18

Thermal demagnetization of the r.m. was carried out by Chamalaun (1963). The directions of n.r.m. when referred to the present horizontal (figure 8) are scattered, mainly in the northeast quadrant with upward inclinations, the distribution of directions being elongate and trending towards the present axial dipole field direction. The n.r.m. directions when referred to the palaeohorizontal still fall in the northeast quadrant but have predominantly downward inclinations (figure 9).

The directions of r.m. after thermal demagnetization at 300 °C are shown relative to the present horizontal in figure 10 and relative to the palaeohorizontal (i.e. to the bedding planes) in figure 11. The former group of directions are more scattered (c.s.d. = 52°) than before cleaning (c.s.d. = 46°) and the mean has moved away from the present geomagnetic field

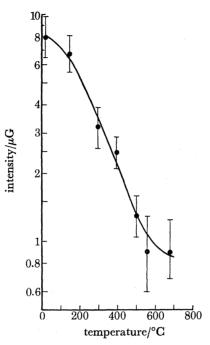


FIGURE 6. Thermal decay of r.m. of Cambrian and Cambro-Ordovician samples. Geometric mean intensity with its standard error at each stage of cleaning plotted on logarithmic scale against temperature.

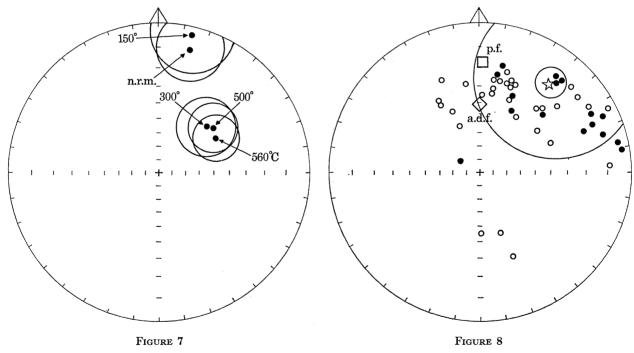


FIGURE 7. Mean direction and standard error circles of r.m. of the second group of Cambrian and Cambro-Ordovician samples subjected to thermal cleaning, shown at each stage.

FIGURE 8. Ordovician samples from Salta and Jujuy. Disk n.r.m. directions plotted with respect to present horizontal. Present field (p.f.) and axial dipole field (a.d.f.) directions are indicated. The mean (star) with standard error and standard deviation circles are also given. One or two disks per sample, 26 samples.

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		TABLE 7. ORDOVICIA	AN SAMPLE DESCRIPT	7. Ordovician sample descriptions from Jujuy Province		
ite no.	sample nos.	location	dip of beds	description	colour	code no.
1	509	Yala–Las Lagunas road	43° at 290°		pale red	5R-6/2
67	510-512	Termes de Cordon	60° at 276°	massive, medium grained sandstone	{dark reddish brown {greyish red	10R-3/4 10R-4/2
က	513-514	At Las Lajitas	30° at 300°	laminated, medium-coarse grained sandstone	greyish red	10R-4/2
4	516-518	by Rio Grande, between sites 3 and 5	25° at 245°	massive, medium grained, sandstone	greyish red	5R-4/2
5	519 - 524	Rio Capillas†	40–60° at 275°	massive mudstone	greyish red	5R-4/2
9	529–530	Volcan–Tumbaya road (road no. 9)	35° at 90° (or 145° at 270° ?)	massive, fine grained, sandstone		
1-	531-537	Quebrada Purmamarca	60° at 295°	massive, fine to very fine grained sandstones	greyish red greyish red	10R-4/2 5R-4/2
œ	525-526	Arrojo el Potrero‡	25° at 270°	massive, very fine grained sandstone	very dusky red dark reddish brown	10R-2/2 10R-3/4

† See Capillas sheet (no. 2566), scale 1:50 000, ref. 406 (N) 900 (E). ‡ See Capillas sheet (no. 2566), scale 1:50 000, ref. 374 (N) 894 (E).

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irecti on (compare figure 10 with figure 8). The grouping of directions referred to the bedding is slightly improved by cleaning (c.s.d. = 28° after cleaning as compared with 32° before) as may be seen in figures 11 and 9. The mean directions and relevant statistical parameters of the four populations of directions are given in table 8.

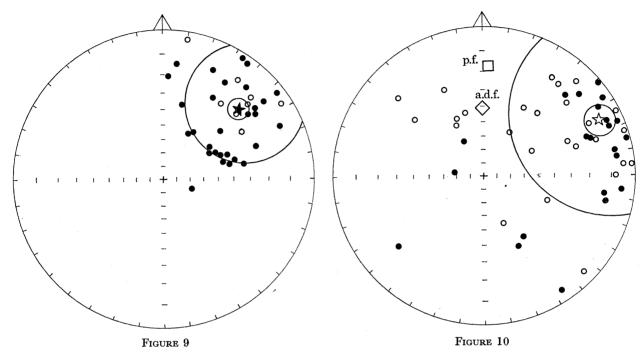


FIGURE 9. Disk n.r.m. directions for the same population of samples as figure 8 plotted with respect to the bedding. FIGURE 10. Disk r.m. directions referred to present horizontal after thermal cleaning at 300 °C. Same population of samples as figure 8.

Table 8. Statistical analysis of r.m. of Ordovician

			di	rection of	r.m./deg	ree
plane of reference	treatment	number of disks	\bigcap	I	δ	$\delta_{ m m}$
present horizontal	n.r.m. 300 °C	47 50	37 63	$-17 \\ -7$	$\begin{array}{c} 46 \\ 52 \end{array}$	6.5 7.5
palaeohorizontal	n.r.m. 300 °C	$\begin{array}{c} 42 \\ 42 \end{array}$	$\begin{array}{c} 46 \\ 52 \end{array}$	$^{+22}_{+53}$	$\begin{array}{c} 32 \\ 28 \end{array}$	$\frac{5}{4.5}$

One or two disks taken from each of 26 samples.

Chamalaun (1963) computed the directions of the low-temperature components which were removed by thermal demagnetization. These may be seen to be alined approximately along the present or axial dipole field directions in figure 12.

Three typical thermal decay curves are shown in figure 13, one for an 'unstable' sample, i.e. one whose r.m. direction changes appreciably during thermal demagnetization and two for 'stable' samples (samples whose magnetization vectors show only a small change). The secondary component in the first sample is removed by treatment below 300 °C, and that is why this temperature was chosen for the bulk demagnetization of the samples.

The south palaeomagnetic pole and relevant statistics are illustrated in figure 18 and table 14.

3.3. Carboniferous-Permian-Triassic

K. M. CREER

The Paganzo formation was sampled in La Rioja Province at sites shown in figure 14. Particulars about the samples are given in table 10 where they have been classified as Permo-Carboniferous or Permo-Triassic according as to whether they come from the bottom part or the top part of the succession.

N.r.m. measurements only have been made, but the more stable samples have been selected by using the criterion that the within-sample circular standard deviation of disk directions should be less than 15°.

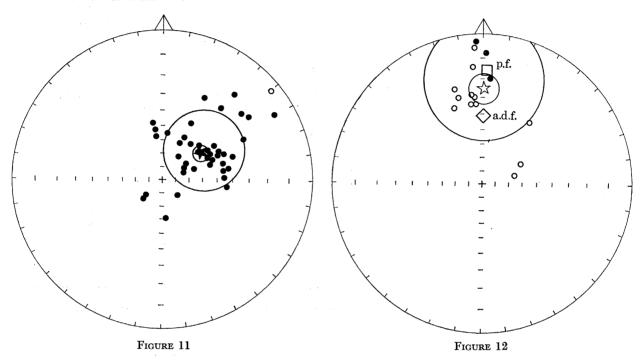


Figure 11. Disk r.m. directions referred to bedding after thermal cleaning at 300° C. Same population of samples

FIGURE 12. Directions of the low-temperature components of r.m. removed by thermal cleaning at 300° C of Ordovician disks for Salta and Jujuy referred to the present horizontal. One disk per sample.

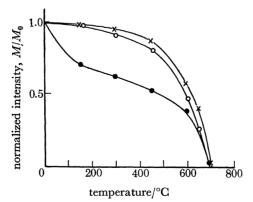


FIGURE 13. Thermal decay curves for Ordovician samples from Salta and Jujuy. The curve showing a sharp decrease below 200 °C is typical of those samples whose r.m. direction changes markedly during cleaning. The other two curves are typical of samples whose r.m. directions show only small changes.

pale reddish brown 10R-5/4: greyish red 10R-4/2

pale reddish brown 10R-5/4

pale reddish brown 10R-5/4

soft m.g. sandstone soft m.g. sandstone soft m.g. sandstone soft m.g. sandstone soft m.g. sandstone

muddy sandstone muddy sandstone

14° at 185° 2° at 100° 17° at 200° 10° at 190°

at approx. km 524 km 525 km 528.5

603 604 605–606

Permo-Triassic

209 809 609

10° at 340° 17° at 350°

km 540km 548km 535

 36° at 0°

km 551

moderate reddish brown 10R-4/6 moderate reddish brown 10R-4/6

greyish red 10R-4/2 greyish red 10R-4/2

moderate reddish brown 10R-5/4: pale reddish pale red 5R-6/2: pale reddish brown 10R-6/6 pale reddish brown 10R-5/4: moderate red pale reddish brown 10R-5/4 colour pale red 10R-6/2 brown 10R-4/6 pale red 5R-6/2 5R - 5/4muddy and c.g. sandstones muddy and c.g. sandstones muddy sandstones muddy sandstones m.-c.g. sandstone lithology marly sandstone 80° to 110° at 240° to 260° 10° to 60° at 34° at 250° 12° at 240° 35° at 130° 12° at 220° dip of beds .08 3 km E. of Puerto Alegre (29° 27' S, 67° 55' W) Zanja de la Vivda Dique los sauces Los Colorados West Huaco locality East Huaco sample nos. 582-586 599 600–601 579 - 581589-593 594-598 Permo-Carboniferous

Table 10. Descriptions of samples and site localities of the La Rioja collection

All sites are on r.n. 40 which is shown in figure 14 and which runs north from San Juan to Chilecito. See sheet 3381 'San Juan', scale 1:1000000 published by the Kilometre references for samples 603-610 are to km posts registering the distance from Mendoza City, Miranda being at km 364, Instituto Geografico Militar (1957), Buenos Aires.

TABLE 11. MEAN DATA FOR LA RIOJA SAMPLES

		$\underset{\widehat{c}^{\mathbf{f}}}{\operatorname{number}}$	dire	direction of n.r.m./degree	.r.m./deg	ree	nos	south palaeopole position	ole posit	ion
age	sample nos.	disks	D	I	ø	$\delta_{ m m}$	lat.	long.	8	4
ermo-Triassic	599-610	22	14	-37	36	∞	S .91	76° S 157° W	36	00
rmo- larboniferous	579–598	15	164	+ 67	16	4	65° S	65° S 44° W	25	9

Table 12. Permo-Carboniferous Salta Province sample descriptions and sites

code no.	10R-6/6 10R-4/2	$10\mathrm{K-}6/2 \ 10\mathrm{R-}6/2 \ 5\mathrm{R-}6/2$	5R-7/4 5R-7/4 5R-5/4
colour	moderate reddish orange greyish red	pale red pale red pale red	moderate pink moderate pink
lithology	soft, friable f.g. sandstone m.g. sandstone	m.g. sandstone	soft, mf.g. sandstone fm.g. sandstone
dip of beds	60° at 260° 60° at 60°	15° at 230°	30° at 340° 60° at 290°
locality	Arrojo del Potrero 48 km from Humahuaca on road to el Aguila	46 km from Humahuaca on road to el Aguila	Quebrada Yacoraite Quebrada de Juella
sample no.	527 554–555	556	569 575–576

TABLE 13. MEAN DATA FOR MESOZOIC FORMATIONS FROM SALTA PROVINCE

ic pole		$\delta_{ m m}$	9
ıet.		40	24
south palaeomagn		long.	$105^{\circ}\mathrm{W}$
nos		lat.	67° S
gree		$\delta_{ m m}$	χÇ
.r.m./deg		9	20
direction of n.r.m./degree		I	+ 59
dire		D	198
number	ot	disks	17

Table 14. Summary of south palaeomagnetic poles deduced from this study

	δ_{m}	14	16	τĊ	9	80	9
south palaeopole	80	09	99	34	25	36	24
	long.	$40^{\circ} \mathrm{W}$	$22^{\circ}\mathrm{W}$		44° W	157° W	$102^{\circ} \mathrm{W}$
	lat.	8	12° N	$11^{\circ}\mathrm{N}$	65° S	8 9	9°
to mortune	measurements	(13†	18‡	428	228	15§	178
	period		Cambrian and Cambro-Ordovician	Ordovician	Permo-Carboniferous	Permo-Triassic	Permo-Triassic
	province	1) 8.14. 2.1 1.	$2 \int$ Salta and Jujuy	3 Salta and Jujuy	4 La Rioja	5 La Rioja	6 Salta and Jujuy

[†] These samples were all thermally cleaned at 560 °C. ‡ These samples were thermally demagnetized at temperatures indicated in tables 5 and 6. § Disks, not samples.

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The Permo-Carboniferous samples are reversely magnetized with a fairly steep inclination, the directions of specimen disks being shown in figure 15. Disk directions from different samples are mixed up, i.e. they do not form clusters about sample-mean directions and this is why they have been plotted rather than sample mean directions.

Table 9. Component of R.M. destroyed by thermal demagnetization compared with PRESENT GEOMAGNETIC FIELD (ORDOVICIAN, §3.2)

	number	direction of r.m./degree				
r.m. or field	disks	D	I	δ	$\delta_{ m m}$	
low temperature component of r.m. referred to present horizontal	15	0	-26	31	8	
present geomagnetic field direction		2	-16	-		
axial dipole field direction		0	-42			

The Permo-Triassic samples are normally magnetized at rather weaker intensities (the geometric mean intensity is 2.1 μ G as compared with 7.5 μ G for the older samples). Disk directions have again been plotted (figure 16).

Mean directions of upper and middle parts (III and II) of the Paganzo formation are given in table 11 with statistical parameters computed using disks and also sample-means as statistical units.

Permo-Triassic formations from the sub-Andean ranges of Jujuy Province have also been studied. The sampling sites are listed in table 12.

The samples with more stable n.r.m. were selected as described for the Paganzo formation. Disk directions are illustrated in figure 17 and the statistics concerning the measurements given in table 13. The palaeomagnetic pole is close to that obtained from the measurements on part III of the Paganzo formation.

South palaeomagnetic pole positions are given in table 14 and have been plotted in figure 18.

4. Discussion of results

4.1. Carrier of remanence

Ore microscope observations have been made on many of the samples studied. All show a finely divided red pigment situated interstitially between quartz and feldspar grains or coating them. After heating in air to 600 °C as is required in thermal demagnetization experiments the pigment in some specimens turns a bluish colour when viewed in reflected light, indicating that the ferric oxides and oxyhydroxides tend to be reduced on heating. In some rocks small cavities (10 to 30 μ m in diameter) appear to be filled with red pigment. Specularite grains typically of 10 to 20 μ m in diameter are observed in most specimens and also in some very small red haematite platelets about 1 μ m in diameter.

Magnetite grains were not observed, although some of the specularite grains appeared to be pseudomorphs of spinel.

The abundance of ferric oxides is consistent with the thermal decay of intensity during demagnetization by cooling in zero magnetic field. The remanent magnetization is not destroyed by heating to 600 °C. Unstable components of remanence are considered to reside in haematite of fine grain size and low blockage temperatures.

4.2. Relevant palaeoclimatic evidence

The evidence concerning upper Palaeozoic glaciations has been recently discussed in a paper by Frakes & Crowell (1969).

It has already been mentioned in §2, that evidence of glaciation has been found in rocks of Mississippian and Pennsylvanian age in northwest Argentina. The palaeomagnetic south pole was only 10°-20° away from this region in the Lower Carboniferous. By the Upper Carboniferous, it was some $40^{\circ}-50^{\circ}$ away (in Kiaman time) and by the Triassic some 60° away. The palaeomagnetic latitudes for the Permo-Carboniferous and Permo-Triassic are consistent with the production of red beds then. The persistence of some indications of glaciation into the Upper Carboniferous may be attributed to Alpine type rather than a polar ice cap.

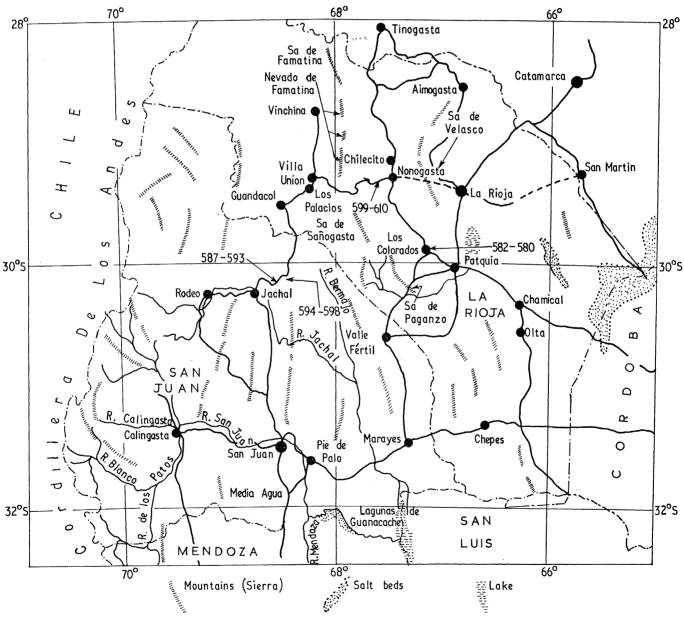


FIGURE 14. Sketch map of area of collection in La Rioja Province. The sampling sites are marked.

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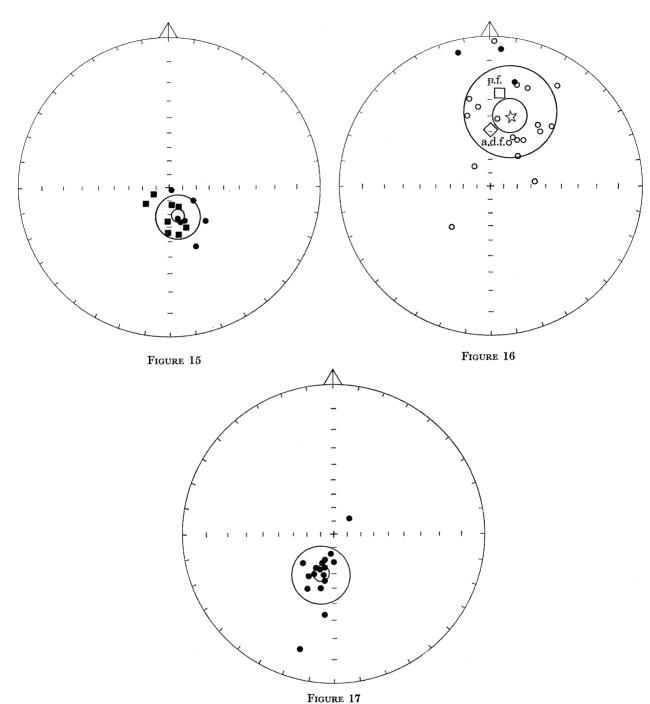


FIGURE 15. N.r.m. directions of specimen disks from Permo-Carboniferous samples from La Rioja Province showing small internal scatter (c.s.d. less than 15°). Mean direction, standard error and standard deviation circles shown. The plane of projection is the palaeo-horizontal. , From east and west Huaco; , from los Colorados and Dique los Sauces.

FIGURE 16. N.r.m. directions of specimen disks from Permo-Triassic samples showing small internal scatter (c.s.d. less than 15°) mean direction, standard error and standard deviation circles shown. The plane of reference is the palaeo-horizontal.

FIGURE 17. N.r.m. directions of specimen disks from Permo-Triassic samples from Jujuy Province showing small internal scatter (c.s.d. less than 15°). Mean direction, standard error and standard deviation circles shown. Plane of reference is the palaeo-horizontal.



FIGURE 18. South palaeomagnetic pole positions with standard error circles for the formations studied. C, CO = Cambrian and Cambro-Ordovician; O = Ordovician; PC = Permo-Carboniferous, PR = Permo-Carbon Triassic. Subscripts refer to entries in table 14, p. 496.

4.3. Consistency with other Palaeozoic palaeomagnetic data from S. America

Palaeomagnetic studies on Bolivian and Brazilian Palaeozoic rock formations lead to the conclusion drawn from the present work that the south palaeomagnetic pole lay to the northeast of Brazil in the lower Palaeozoic and that it moved southwards down the eastern side of S. America to a position in the S. Atlantic east of the River Plate approximately in the Upper Palaeozoic.

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as part of their Ph.D. projects. The Upper Palaeozoic collection has since been incorporated into a new and much larger collection made in 1968-9 by Ing. D. Valencio, Mr F. Vilas, Mr B. J. J. Embleton and myself and which is being studied by B.J.J.E. as part of his Ph.D. research project.

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